

Co-seismic strain-steps associated with the 2004 off the Kii peninsula earthquakes—Observed with Ishii-type borehole strainmeters and quartz-tube extensometers

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The 2004 off the Kii peninsula earthquakes (Mj7.1 and Mj7.4) occurred at the Nankai trough, on 5 September 2004. Clear strain-steps associated with these earthquakes were observed with Ishii-type borehole strainmeters and quartz-tube extensometers in the Tokai and Kinki districts. We investigated the spatial and depth distribution of the observed principal strain changes and compared the observed strain-steps and theoretical calculations at all observatories. The following results were obtained: the observed strain-steps at all observatories are generally consistent with the polarities of the theoretical values, and the observed strain-step increases with depth at the same place. As for the Togari site, the following relationships are obtained: the strain-step and the tidal strains increase with depth and increasing of the modulus of elasticity, namely, hardness of rock. We consider that the geological structure around the observatory may cause a modification of the strain field.

Key words: The 2004 off the Kii peninsula earthquakes, strain-step, principal strain, Ishii-type borehole strainmeter, quartz-tube extensometer, modification of strain field.

1. Introduction

Strain-steps associated with large earthquakes have been reported during the last few decades. For example, Kasahara (1974) and Okada (1975) investigated the space distribution of the co-seismic strain-steps observed with extensometers which are widely distributed throughout Japan, and compared the co-seismic strain-steps with the theoretical values from a dislocation model of the earthquakes, showing many strain steps which agree with the polarity of the theoretical value and some strain steps which do not agree. They suggested the mechanical instability of the observation equipment, groundwater flow change, etc., as possible causes. However, the causes are inconclusive.

During the last 10 years, in the Tokai and Kinki districts of central Japan, several crustal movement observatories were built by the Tono Research Institute of Earthquake Science (TRIES) and Geological Survey of Japan (GSJ), and stable continuous observations have been performed with a 20 bit A/D resolution, far superior to that of the previous resolution. In the Tokai district, Nagoya University has been performing continuous crustal movement observations for the last 30 years. Under these circumstances, the 2004 off the Kii peninsula earthquakes (Mj7.1 foreshock and Mj7.4 main shock) occurred at the Nankai trough at 19:07 and 23:57 (JST), 5 September 2004, respectively (Fig. 1).

Clear strain-steps associated with these earthquakes were observed at each observatory. In this study, we investigate the spatial and depth distribution of the observed strain-steps and compare the observed steps with the theoretical values from a dislocation model of the earthquakes, and attempt to clarify the differences of the strain field.

2. Observation

We analyzed the crustal strain data of the observatories shown in Fig. 1 and Table 1. The Byobu-san borehole observatory (BYB) is the deepest multi-component continuous monitoring of crustal activities in the world (Ishii *et al.*, 2003). The Togari borehole observatory consists of two boreholes with depths of 350 m (TGR350) and 165 m (TGR165). These two boreholes were drilled about 3 m apart. The instruments were installed in the granite bedrock all three borehole observatories, which belong to TRIES. At the Mizunami vault-type observatory (NAMZ) is located in the Tertiary sedimentary rocks. The Asahi (NAAS), the Inabu (NAIB) and the Toyohashi (NATY) are vault-type observatories. NAAS and NAIB are located in the granite. NATY is located in the pre-Tertiary sedimentary rocks. These four vault-type observatories belong to Nagoya University. At the Toyohashi borehole observatory (TYH1) and the Ohara borehole observatory (OHR), belonging to GSJ, the instruments were installed in the pre-Tertiary sedimentary rocks and pre-Tertiary volcanic rocks, respectively. At these borehole and vault-type observatories, crustal strains were measured by Ishii-type borehole strainmeters (Ishii *et*

Table 1. List of the observatories.

Station	Latitude (deg)	Longitude (deg)	Altitude (m)	Depth (m)	Geology	Density (g/cm ³)	Vp (km/s)	Vs (km/s)	Shear modulus (GPa)	Bulk modulus (GPa)	Young modulus (GPa)	Poisson ratio
BYB	35.3643	137.3039	333	1020	G	2.62	4.72	2.71	19.24	32.71	48.26	0.25
Togari site												
TGR350	35.3745	137.2378	190	350	G	2.59	5.18	2.57	17.11	46.69	45.73	0.34
TGR165	35.3745	137.2378	190	165	G	2.54	4.23	1.64	6.83	36.34	19.29	0.41
NAMZ	35.3755	137.2388	195	20	T	1.83	2.08	1.00	1.83	5.48	4.93	0.35
NAAS	35.2249	137.3620	200	50	G							
NAIB	35.2015	137.5337	700	50	G							
Toyohashi site												
TYH1	34.7643	137.4667	77	250	pT	2.69	4.58	2.53	17.33	33.69	44.39	0.28
NATY	34.7652	137.4667	76	50	pT							
OHR	35.1105	135.8193	217	278	G							

G: Granite, T: Tertiary sedimentary rocks, pT: pre-Tertiary sedimentary rocks.

Table 2. Comparison of the observed strain-steps (Obs) and theoretical calculations (M-1 to M-4) for the foreshock (upper table), and the main shock (lower table). α is a modification coefficient. Unit of observed strain-step (Obs) and theoretical calculations are nano-strain. Unit of the direction of the maximum principal strain is degrees; clockwise from the north are positive.

Foreshock		Maximum Principal strain						Minimum Principal strain						Direction of Maximum Principal strain				
Mj7.1	Obs	M-1	M-2	M-3	M-4	α	Obs	M-1	M-2	M-3	M-4	α	Obs	M-1	M-2	M-3	M-4	
BYB	116.0	52.1	50.6	53.3	74.9	2.3	-155.5	-18.4	-17.5	-17.8	-25.6	8.9	15.5	14.4	13.4	13.2	12.6	
Togari site																		
TGR350	88.3	52.9	51.1	53.8	75.7	1.7	-42.1	-18.4	-17.6	-17.9	-25.6	2.4	21.7	12.8	11.7	11.5	10.9	
TGR165	34.4	52.9	51.1	53.8	75.7	0.7	-34.1	-18.4	-17.5	-17.9	-25.6	1.9	6.1	12.8	11.7	11.5	10.9	
NAMZ	37.9	53.4	51.5	54.2	76.3	0.7	-16.6	-18.5	-17.6	-17.9	-25.7	0.9	28.0	12.1	11.1	10.9	10.3	
NAAS	33.5	59.8	58.4	62.0	87.2	0.6	-11.7	-21.5	-20.5	-21.0	-30.2	0.6	10.3	16.8	15.7	15.6	14.9	
NAIB	28.6	55.6	55.0	58.5	82.3	0.5	-7.8	-20.8	-19.9	-20.3	-29.3	0.4	17.5	21.2	20.1	20.1	19.4	
Toyohashi site																		
TYH1	83.6	98.1	97.7	106.5	151.8	0.9	-52.6	-38.5	-37.0	-38.7	-56.5	1.4	30.3	24.5	23.3	23.4	22.3	
NATY	50.9	99.1	98.6	107.6	153.3	0.5	-22.5	-38.8	-37.3	-39.0	-56.9	0.6	19.5	24.3	23.1	23.2	22.0	
OHR	77.3	68.4	60.3	61.4	89.4	1.3	-19.5	-23.0	-21.1	-21.5	-30.5	0.9	-39.9	-23.0	-24.2	-25.1	-25.1	
Main shock		Maximum Principal strain						Minimum Principal strain						Direction of Maximum Principal strain				
Mj7.4	Obs	M-1	M-2	M-3	M-4	α	Obs	M-1	M-2	M-3	M-4	α	Obs	M-1	M-2	M-3	M-4	
BYB	325.1	96.3	104.3	100.5	79.6	3.1	-406.5	-27.6	-25.8	-25.7	-20.1	15.8	13.3	3.3	4.4	5.2	2.5	
Togari site																		
TGR350	516.9	95.1	103.3	99.8	78.4	5.0	-166.6	-27.1	-25.5	-25.3	-19.7	6.5	33.5	1.4	2.5	3.3	0.4	
TGR165	110.5	95.1	103.3	99.8	78.4	1.1	-86.2	-27.1	-25.4	-25.3	-19.7	3.4	4.3	1.4	2.5	3.3	0.4	
NAMZ	42.5	94.9	103.3	99.8	78.2	0.4	-75.0	-27.0	-25.4	-25.3	-19.6	3.0	39.4	0.7	1.7	2.5	-0.4	
NAAS	81.2	115.5	125.0	120.1	96.9	0.6	-24.2	-33.3	-31.0	-30.7	-24.6	0.8	3.0	5.5	6.8	7.7	4.9	
NAIB	72.2	115.9	123.3	117.7	97.4	0.6	-7.1	-33.6	-30.9	-30.5	-24.8	0.2	15.6	10.7	12.3	13.2	10.6	
Toyohashi site																		
TYH1	379.1	220.4	239.5	226.1	194.5	1.6	-179.5	-67.8	-62.8	-61.2	-53.4	2.9	28.6	12.0	14.2	15.4	12.4	
NATY	160.9	221.9	241.4	228.0	196.0	0.7	-34.0	-68.2	-63.1	-61.5	-53.7	0.5	10.4	11.7	13.9	15.2	12.1	
OHR	189.0	51.4	59.0	59.0	34.5	3.2	-65.3	-22.9	-21.8	-22.5	-18.2	3.0	-59.1	-44.8	-42.9	-42.3	-55.5	

with cut-off frequency of 30 min; Asai *et al.*, 2001) before extracting hourly data from the original data [2]. Using the tidal analysis program BAYTAP-G (Tamura *et al.*, 1991), we calculated the solid Earth tide component and the atmospheric pressure response of each data; they were then removed from the data and the residual component was obtained [3]. Assuming plane strain, the principal strain was calculated from the residuals of the three independent directions and the size of step and the direction of maximum principal strain associated with the foreshock and the main shock were calculated after subtracting a linear trend fit to 115 h of data preceding the foreshock so as to minimize the strain variation before the earthquake (Obs in Table 2). Figure 2 shows the hourly maximum principal strain of all observatories from 1 to 10 September as an example.

The theoretical principal strain associated with the foreshock and the main shock at each observatory were calculated by the four fault models of the Geographical Survey Institute (GSI) (M-1 to M-4 in Table 3; Geographical Survey Institute, 2004) which are inferred from the co-seismic horizontal movements by GEONET (GPS Earth Observation Network; Miyazaki *et al.*, 1997). The Earth's crust was approximated as an isotropic homogeneous semi-infinite elastic medium. The altitudes of vault-type obser-

vatories were set at sea-level, and the installation depths of borehole equipment were set below sea-level. Using these model parameters with the MICAP-G program (Naito and Yoshikawa, 1999) based on Okada (1992), we calculated the principal strain and direction of the maximum principal strain at each observatory (M-1 to M-4 in Table 2).

4. Discussion and Conclusion

As space is limited, we have only focused on the differences of the amplitude of principal strain. The theoretical values of the principal strains in each observatory are almost the same, while the values from GSI's model M-4 are slightly different compared to the other models (see Table 2). As for GSI's models M-2 and M-3, these models are consistent with the aftershock distribution and the fault model on the basis of the displacement field obtained by a campaign survey of GPS in the southern Kii peninsula (Hashimoto *et al.*, 2005). Therefore, comparison of the theoretical values, which were calculated from GSI's model M-2 of the foreshock, and those observed are shown in Fig. 3(a) as an example. Similar plots of observed strain-steps for the main shock are shown in Fig. 3(b). The theoretical values at different depths at the Togari site (TGR165, TGR350 and NAMZ) and the Toyohashi site (TYH1 and

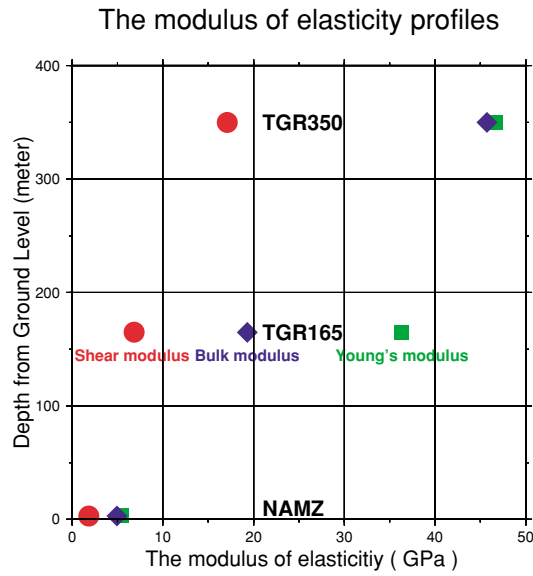


Fig. 4. The modulus of elasticity at different depths at the Togari site.

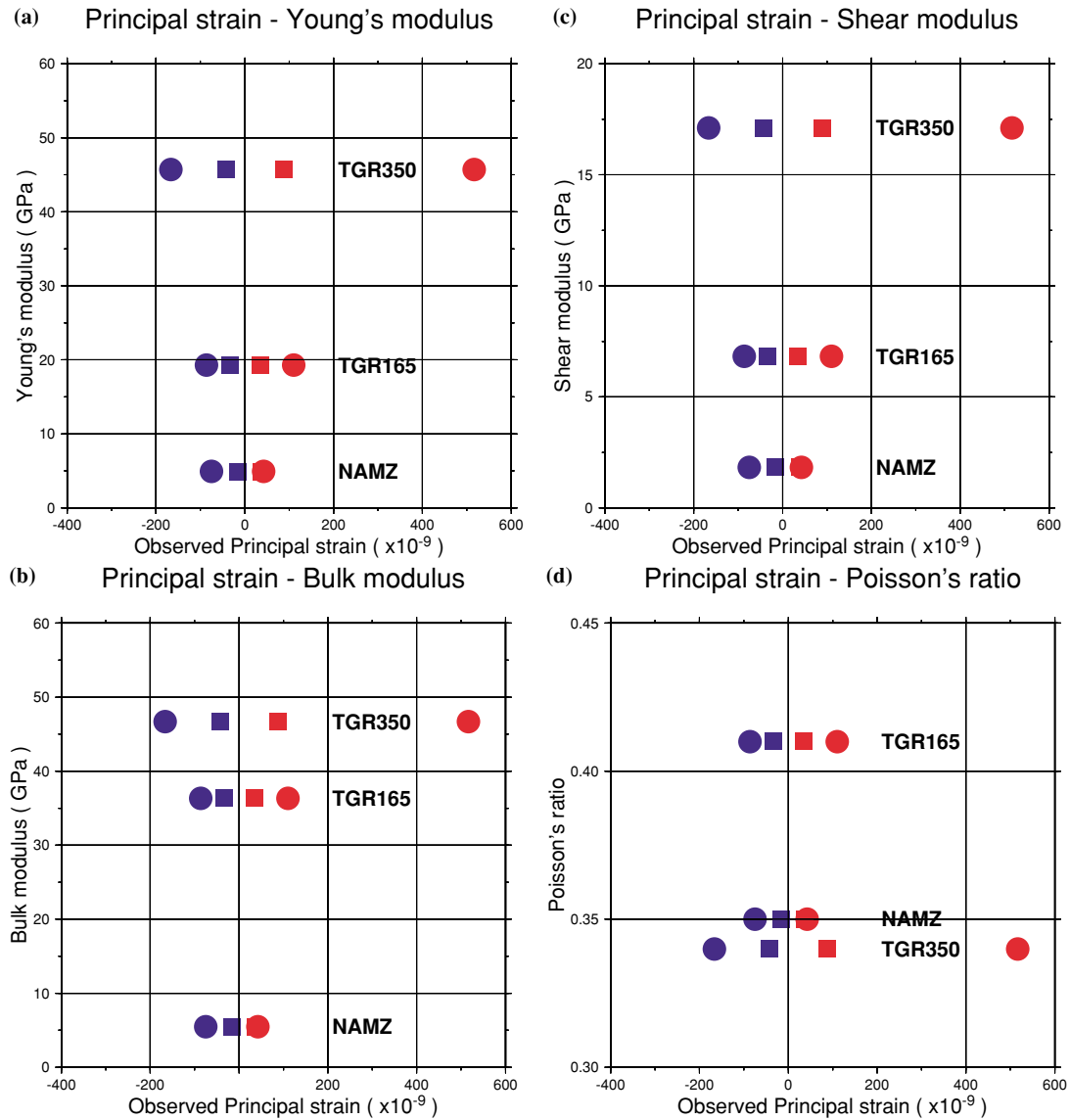


Fig. 5. Comparison of the observed strain-step of maximum (red marks) and minimum (blue marks) principal strain for the foreshock (squares) and the main shock (circles) with Young's modulus (a), bulk modulus (b), shear modulus (c), and Poisson's ratio (d), respectively.

Table 3. GSI's (2004) fault model for the foreshock (upper table) and the main shock (lower table).

Foreshock	Latitude	Longitude	Length	Width	Depth	Strike	Dip	Rake	Slip	Mw
Mj7.1	(deg)	(deg)	(km)	(km)	(km)	(deg)	(deg)	(deg)	(m)	
M-1	32.97	136.95	30.1	19.8	9.3	267	47	98	4.27	7.2
M-2	32.97	136.95	30.0	19.7	11.3	267	47	92	4.00	7.2
M-3	32.97	136.94	30.1	19.3	12	280	41	108	3.83	7.1
M-4	33.03	136.67	30.1	19.9	11.2	71	58	76	5.59	7.3
Main shock	Latitude	Longitude	Length	Width	Depth	Strike	Dip	Rake	Slip	Mw
Mj7.4	(deg)	(deg)	(km)	(km)	(km)	(deg)	(deg)	(deg)	(m)	
M-1	33.18	137.25	34.4	21.5	15.3	245	62	50	5.1	7.3
M-2	33.11	137.29	34.8	18.1	16.9	245	54	62	3.44	7.3
	33.19	137.21	28.3	24.1	5.8	321	85	141	2.84	
M-3	33.11	137.32	38	25.1	17.1	265	52	88	2.28	7.3
	33.16	137.18	32.5	15.4	7.5	323	90	150	3.39	
M-4	33.41	136.67	30.1	19.9	11.2	71	58	76	5.59	7.2

Table 4. Tidal analysis results of maximum and minimum principal strains concerning the Togari site and Toyohashi site, for one month in September 2004. We focused on M2 and O1 constituents because the disturbance of meteorological phenomena does not have much affect much in these constituents. The unit of Amplitude (Amp.) is nano-strain. Phase is the phase shift with respect to local potential; positive value indicates phase lead. RMSE is the root mean square error.

	Maximum Principal strain								Minimum Principal strain							
	M2				O1				M2				O1			
	Amp. (RMSE)	Phase (RMSE)	Amp. (RMSE)	Phase (RMSE)	Amp. (RMSE)	Phase (RMSE)	Amp. (RMSE)	Phase (RMSE)	Amp. (RMSE)	Phase (RMSE)	Amp. (RMSE)	Phase (RMSE)	Amp. (RMSE)	Phase (RMSE)	Amp. (RMSE)	Phase (RMSE)
Togari site																
TGR350	11.4	0.2	-16.3	1.0	4.6	0.7	43.9	9.0	11.0	0.3	15.6	1.3	9.5	0.8	-31.1	4.6
TGR165	6.9	0.4	-23.0	2.9	5.4	1.3	27.4	13.3	5.1	0.4	30.7	4.0	4.2	1.2	-46.2	17.0
NAMZ	6.8	0.1	-21.2	1.1	5.0	0.4	20.7	5.1	5.5	0.1	8.3	1.2	3.7	0.4	26.8	5.9
Toyohashi site																
TYH1	4.7	0.5	-15.1	6.4	2.5	0.5	-12.3	11.8	4.1	0.6	8.7	7.8	2.3	0.5	2.7	13.4
NATY	4.3	0.1	-35.0	1.4	1.6	0.4	-10.0	13.0	2.4	0.1	9.5	2.0	3.1	0.3	-10.1	5.4

NATY) agreed in the nano-strain order (see Table 2). This shows that in these regions, which are more than 180 km from the hypocenter, there is no difference in the depth for the principal strain changes required in the isotropic homogeneous semi-infinite elastic medium model.

The observed strain-steps for the foreshock and the main shock at all observatories are generally consistent with the calculated polarities (Table 2, Figs. 3(a) and 3(b)). These agreements are surprising when compared to earlier reports that some observed strain-steps do not agree with the polarity of the calculated values (e.g. Kasahara, 1974; Okada, 1975), although they were directly compared to the strain-steps of each component. For the NAMZ, NAAS, NAIB, and NATY in the vault-type observatories, a device that suppresses the oscillations of extensometers during earthquakes (Yamauchi *et al.*, 1975) appears to be a major factor in obtaining stable and sensible directional data.

At the Togari site (TGR350, TGR165 and NAMZ) and the Toyohashi site (TYH1 and NATY), there are clear differences in the observed strain-steps at different depths for the foreshock and the main shock. To make these differences clear, we obtained the modification coefficients α ($\alpha > 0$) which are expressed as follows:

$$\alpha = \varepsilon(\text{obs})/\varepsilon(\text{theo})$$

where $\varepsilon(\text{obs})$ is the observed strain and $\varepsilon(\text{theo})$ is the theoretical strain which is calculated from the model. The coefficients α for each observatory using GSI's model M-2 are shown in Table 2. At these sites, the coefficients α for the borehole observatory are larger than those for the vault-type (Table 2, Figs. 3(a) and 3(b)). Let us consider

the coefficients α from the differences in the geological environments at these observatories. As mentioned in Section 2, TGR350 and TGR165 are located at different depths in granite bedrock. NAMZ is located in Tertiary sedimentary rocks. TYH1 and NATY are located at different depths in pre-Tertiary sedimentary rocks consisting of inclined layers of sandstone, chert and shale. Thus, we investigated the relation between the observed strain-steps and the geological environments using the physical properties of the rocks at each depth at the Togari site. As for the Toyohashi site, no investigation has been carried out, because we have no definite information on the physical properties of the rocks in NATY. The bulk modulus, the shear modulus, Young's modulus and Poisson's ratio at each depth were obtained from the P -wave and S -wave velocity and the density of rock at each depth, which were obtained from well loggings or examination of drilling core samples. Figure 4 shows the modulus of elasticity for each depth at the Togari site, and the modulus of elasticity increases with the depth. Figure 5(a) shows that the observed strain-step increases with Young's modulus. A similar plot for the bulk modulus, and shear modulus is shown in Figs. 5(b) and 5(c), respectively, excluding Poisson's ratio (Fig. 5(e)). Similar tendencies are observed in the tidal amplitudes, which obtained the tidal analysis of principal strain for a month in September 2004 (see Table 4 and Asai *et al.*, 2001). At the Togari site, the coefficients α increases with depth and an increasing of the modulus of elasticity, namely, hardness of rock. This is strange because generally the strain in the rock decreases with the hardness.

As for geological structures, between TGR350, TGR165

and NAMZ, there exists a sub-vertical fault with a NNW strike which was recognized geologically (JNC, 2003). TGR165 is located near the boundary between the Tertiary layers and granite bedrock with a depth of about 100 m, where there are many cracks. There are also several confined aquifers at a depth of 200–360 m in the granite bedrock near the installation depth for TGR350 and TGR165, consisting of a joint system or cracks (JNC, 2003). We consider that these differences of the geological structure around the observatory, namely the heterogeneity of rocks, may cause a modification of the strain field. A detailed investigation of the modification is the subject for future study.

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